

Improving our understanding of carbon cycle, drought and fire dynamics during the 21st century (and beyond!)

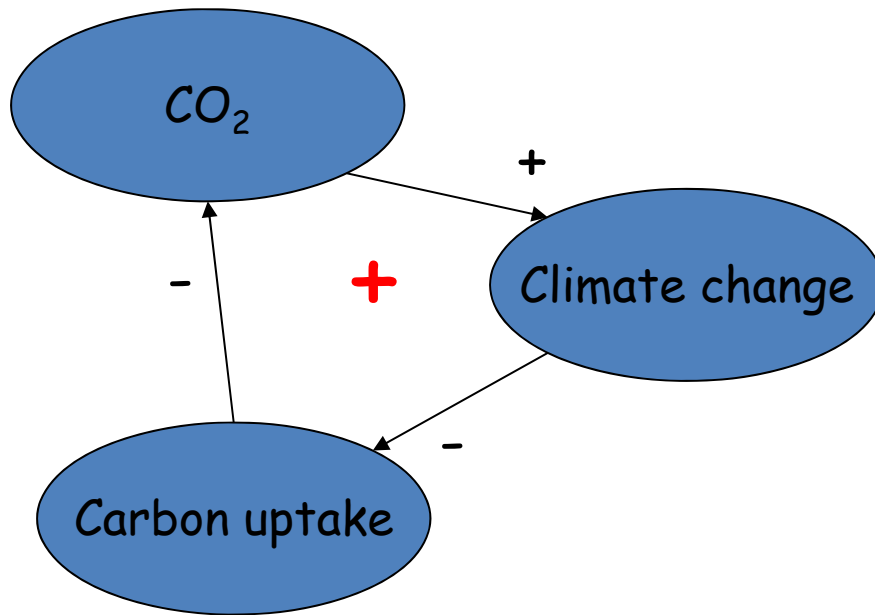
James Randerson
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Earth System Science
UC Irvine

28 October 2015



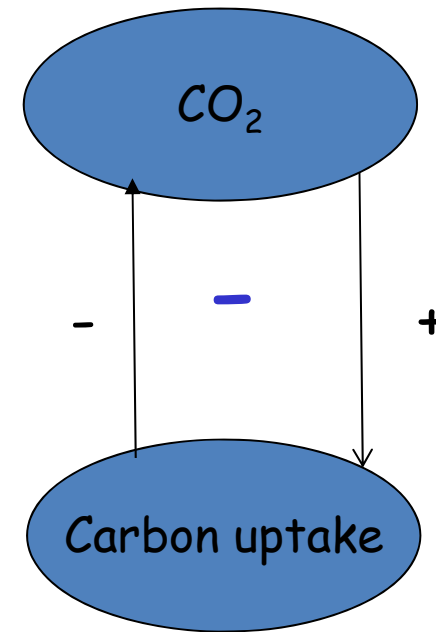
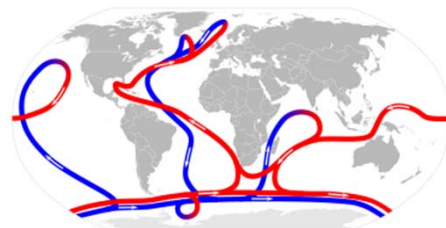
U.S. Dept. of Energy
Biological and Environmental Research
Advisory Committee Meeting

Two types of carbon feedback loops influence the temporal evolution of atmospheric CO₂



Climate-carbon feedback

γ



Concentration-carbon feedback

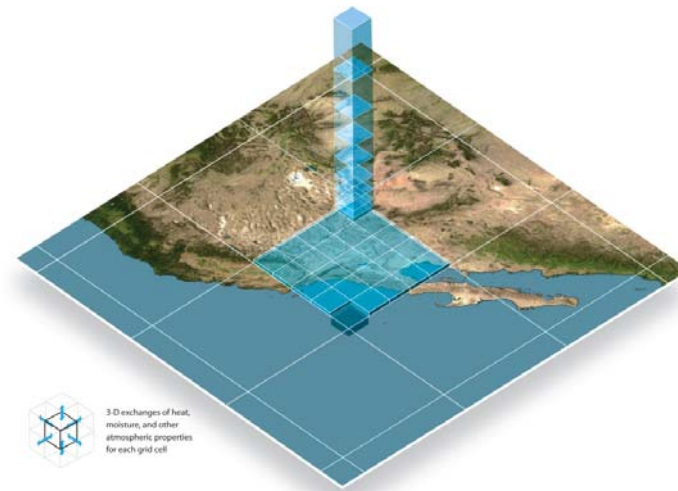
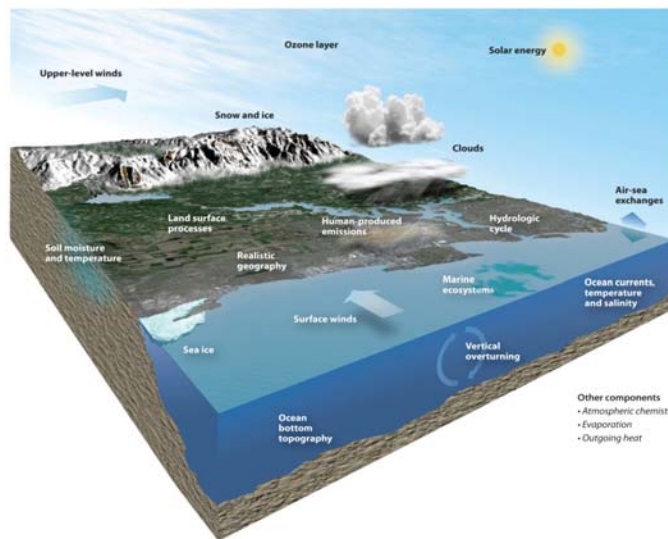
β



Science questions:

- How are ocean and land contributions to the climate-carbon feedback likely to evolve over time?
- How can we use isotope observations to reduce uncertainties in future projections of the soil carbon sink?
- How will climate change influence drought and fire dynamics?

The Community Earth System Model



Graphic credit: UCAR

What are important climate-carbon processes and feedbacks?

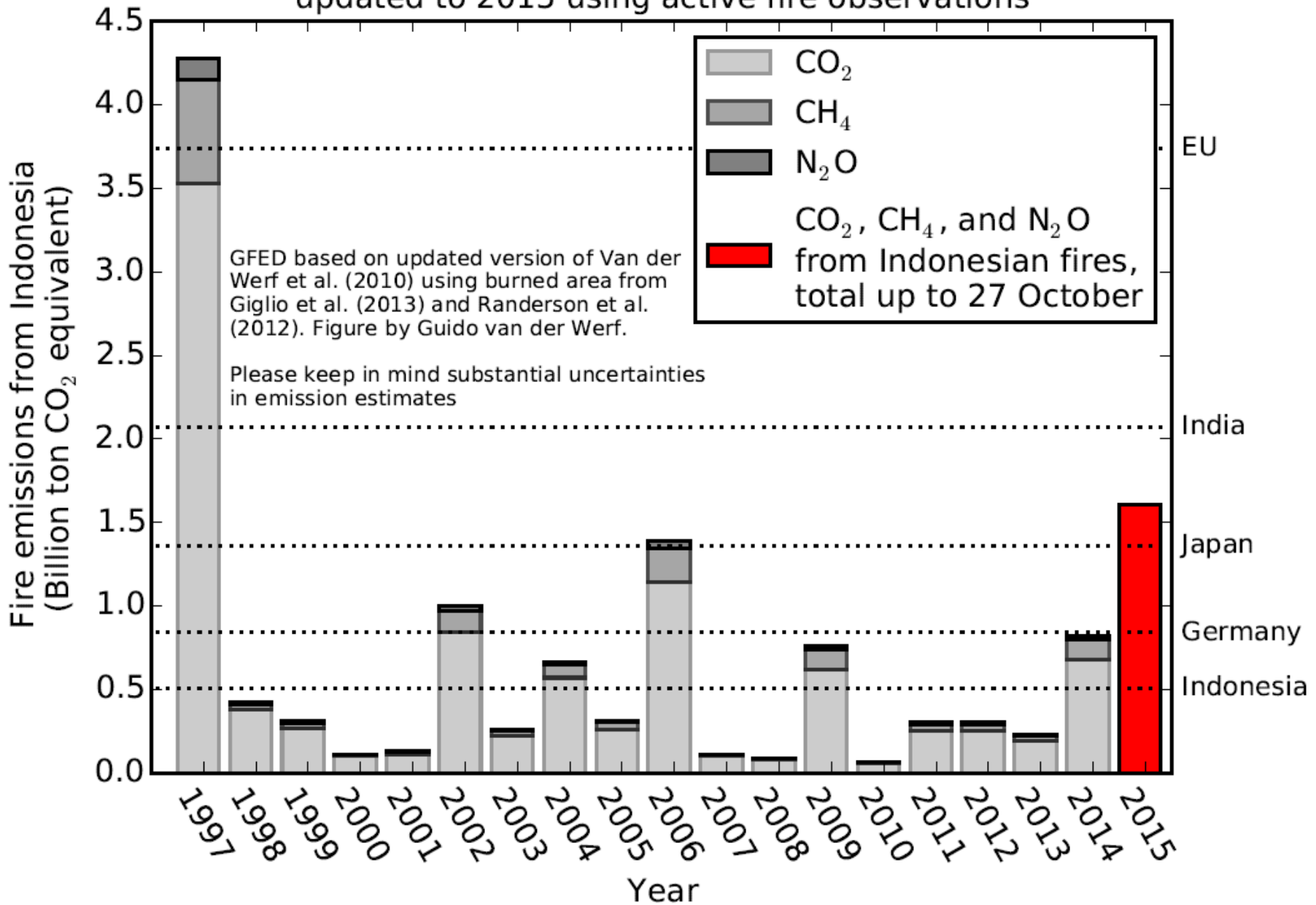
Processes in CESM1(BGC):

- Ocean:
 - Increasing stratification with warming
 - Dissolved inorganic carbon sensitivity to temperature
 - Biological pump responses to stratification
- Land:
 - Drought & temperature effects on primary production
 - Soil decomposition increases in response to temperature
 - Response of fires to changes in fuels and drought
 - Land use change

Not yet in most CMIP ESMs:

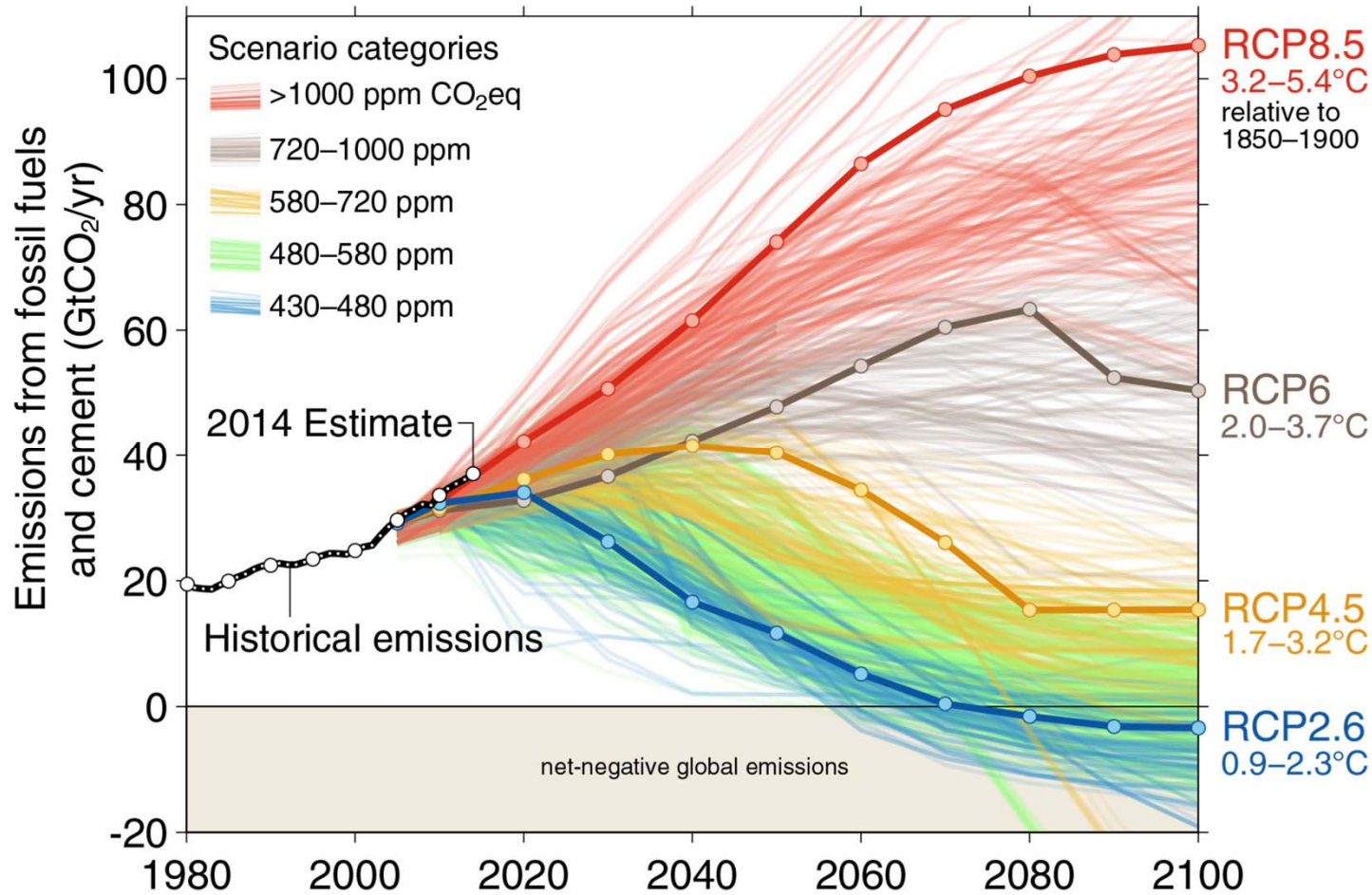
- Species range shifts
- Phosphorus limits on land carbon uptake (integration underway into ACME)
- Permafrost dynamics (now in CLM4.5)
- Peatlands
- Fires
- Insect-driven mortality
- Drought effects on tree mortality
- Climate effects on land use change

Global Fire Emissions Database (GFED) updated to 2015 using active fire observations



Fossil fuel CO₂ emissions for various countries in the year 2013 based on the EDGAR database

Experimental design: All three simulations have prescribed atm. CO₂ from RCP8.5



CESM1(BGC) experimental design

Simulation	Short name	Description
Fully coupled	Full	CO ₂ and other atmospheric anthropogenic drivers influence radiative transfer, biogeochemistry responds to CO ₂ increases
No CO ₂ radiative forcing	No CO ₂ forcing	Non-CO ₂ anthropogenic drivers influence radiative transfer, biogeochemistry responds to CO ₂ increases
No anthropogenic radiative forcing from greenhouse gases or aerosols	No anthro. forcing	No atmospheric anthropogenic climate change, biogeochemistry responds to CO ₂ increases

Validation:

Lindsay et al. (2014), Moore et al. (2013), Long et al. (2013), Keppel-Aleks et al. (2013)

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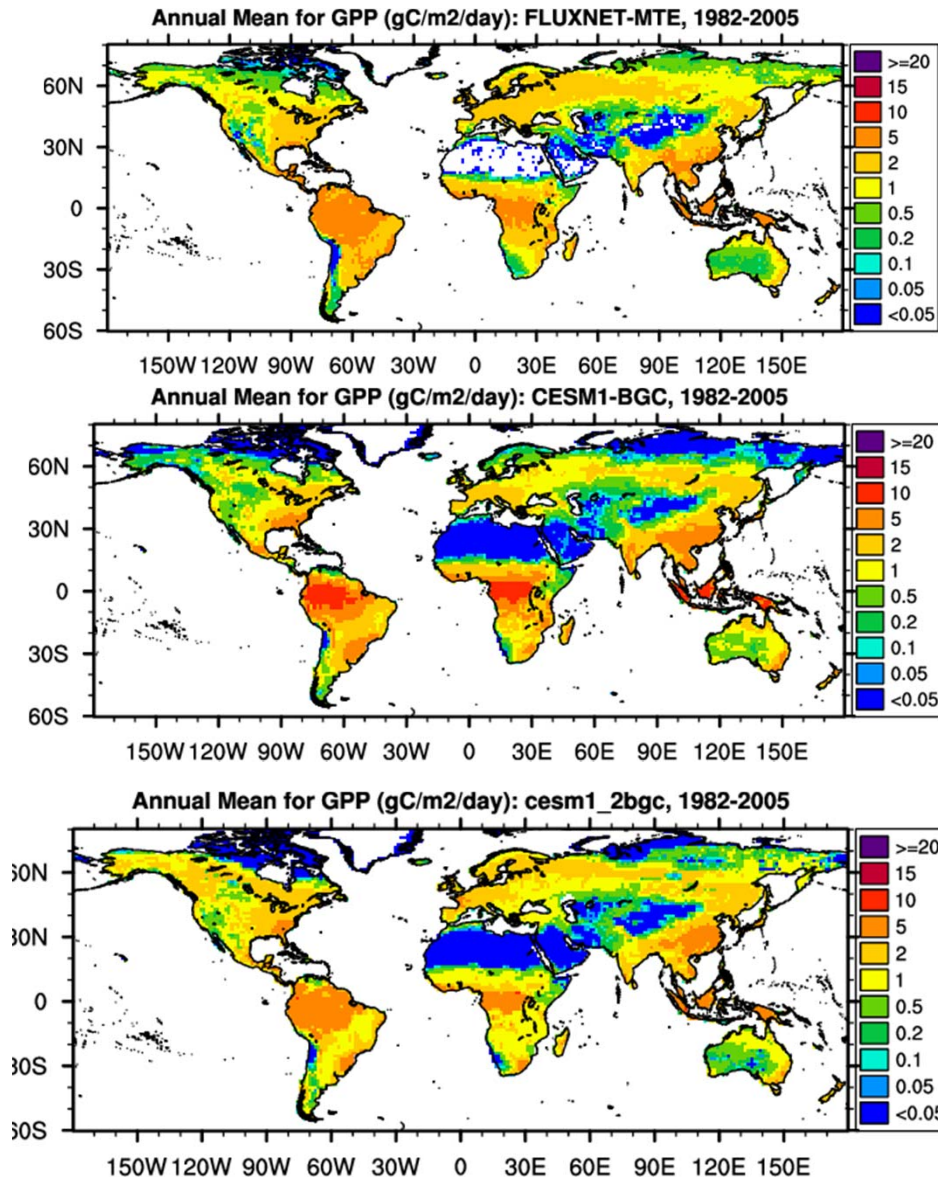
Lindsay et al. (2014), Keppel-Aleks et al. (2013), Moore et al. (2013), Long et al. (2013)

CESM1(BGC) experimental design

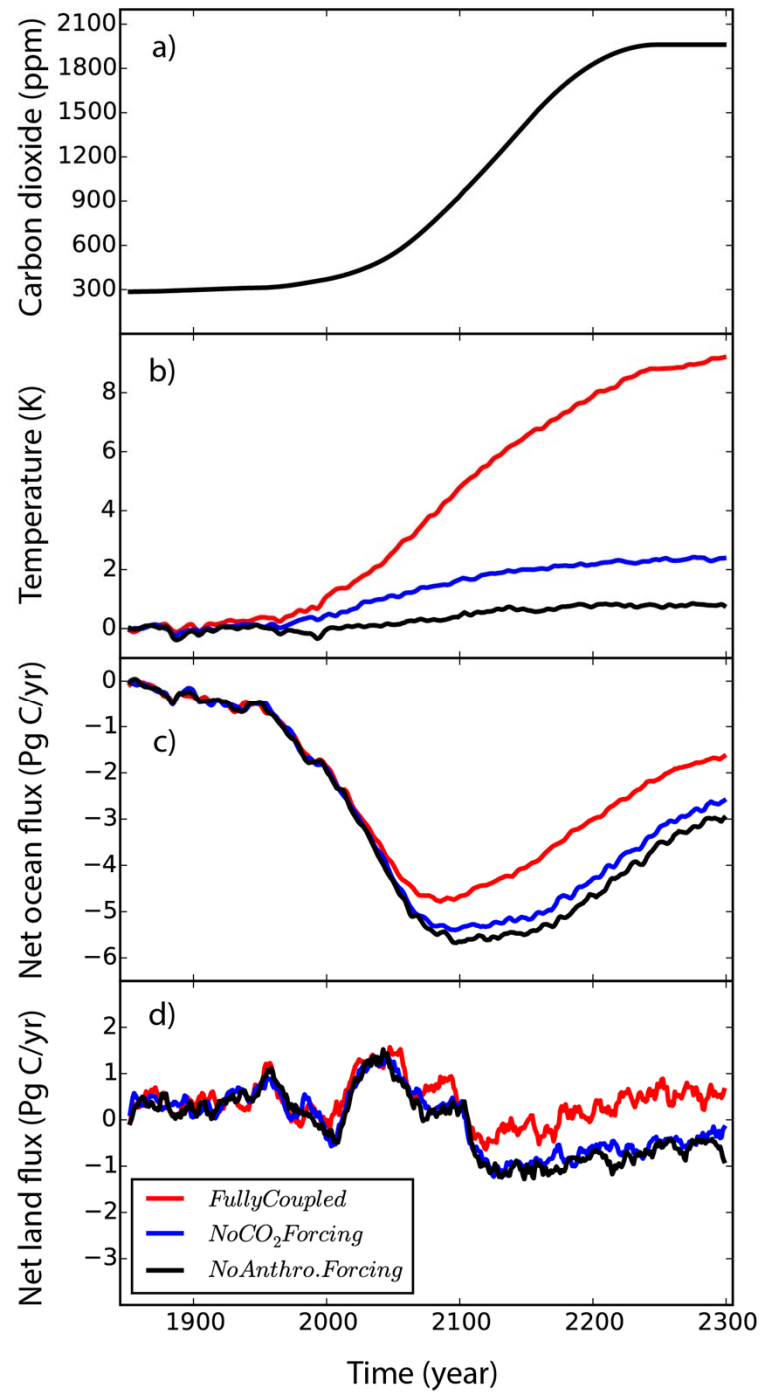
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No anthropogenic radiative forcing from greenhouse gases or aerosols	No anthro. forcing	No atmospheric anthropogenic climate change, biogeochemistry responds to CO ₂ increases

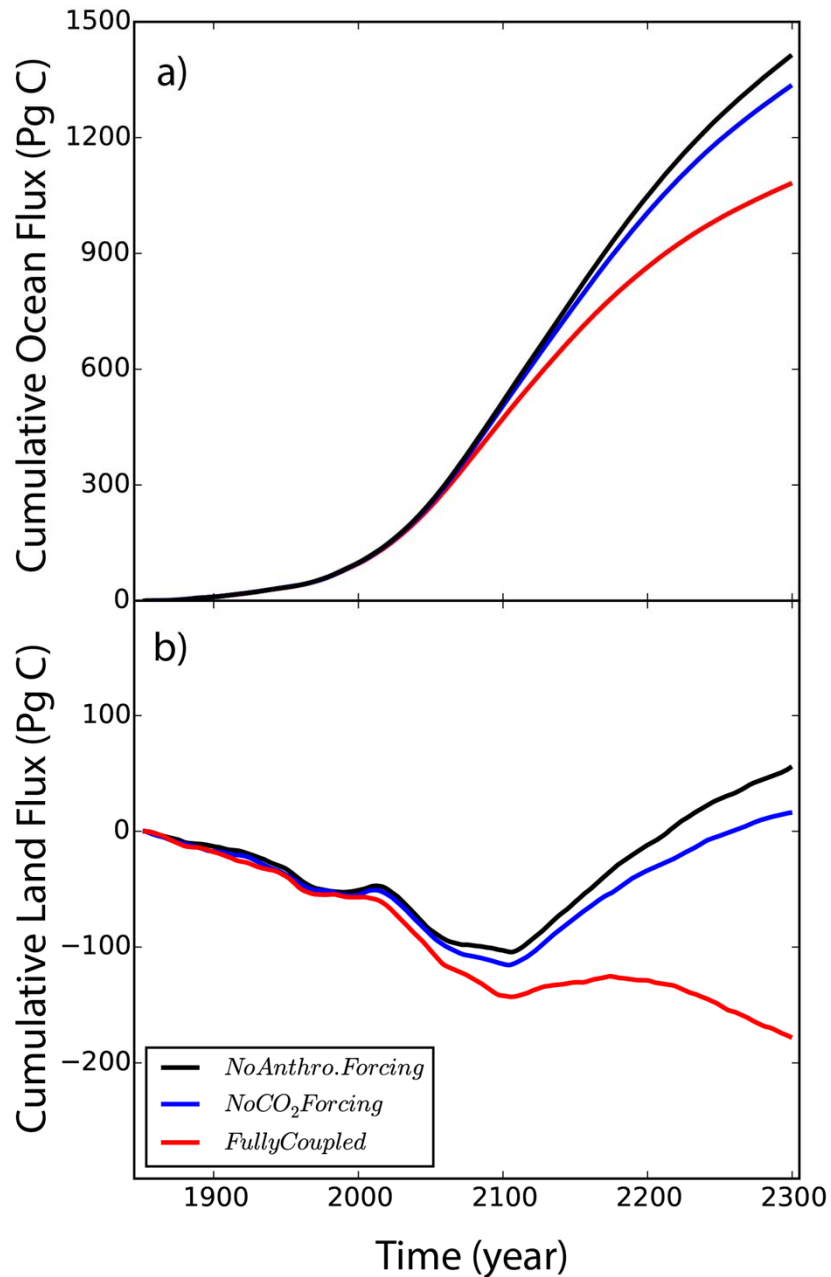
Lindsay et al. (2014), Keppel-Aleks et al. (2013), Moore et al. (2013), Long et al. (2013)

Validation of carbon cycle processes in CESM with the International Land Model Benchmarking System



Process and number of variables	CESM 1	CESM 1.2
Carbon cycle and ecosystems (9)	0.54	0.61
Hydrological cycle (3)	0.66	0.77
Radiation and energy (7)	0.80	0.80
Forcing variables (4)	0.83	0.84
Variable to variable relationships (10)	0.67	0.71
Overall score:	0.65	0.69





Climate-carbon gain computed from compatible fossil fuel emissions (E) from fully coupled and no CO_2 forcing simulations

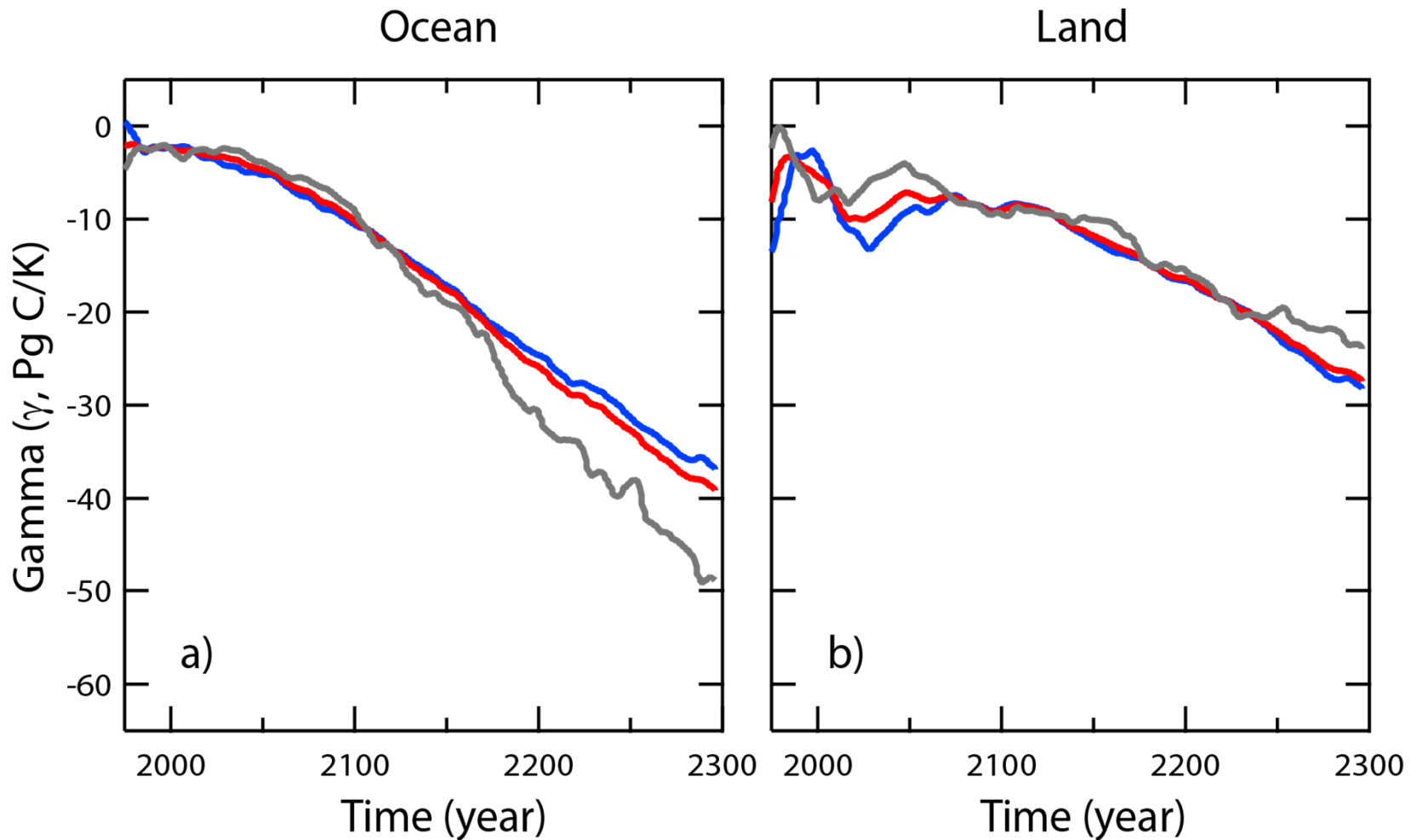
$$g = \frac{E_{noCO_2} - E_{FC}}{E_{noCO_2}}$$

Climate-carbon feedback parameters

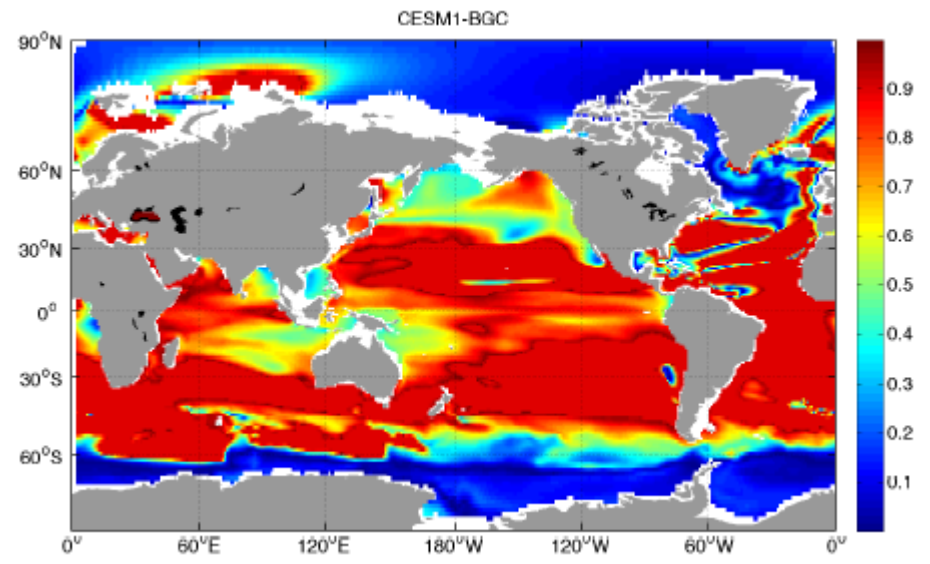
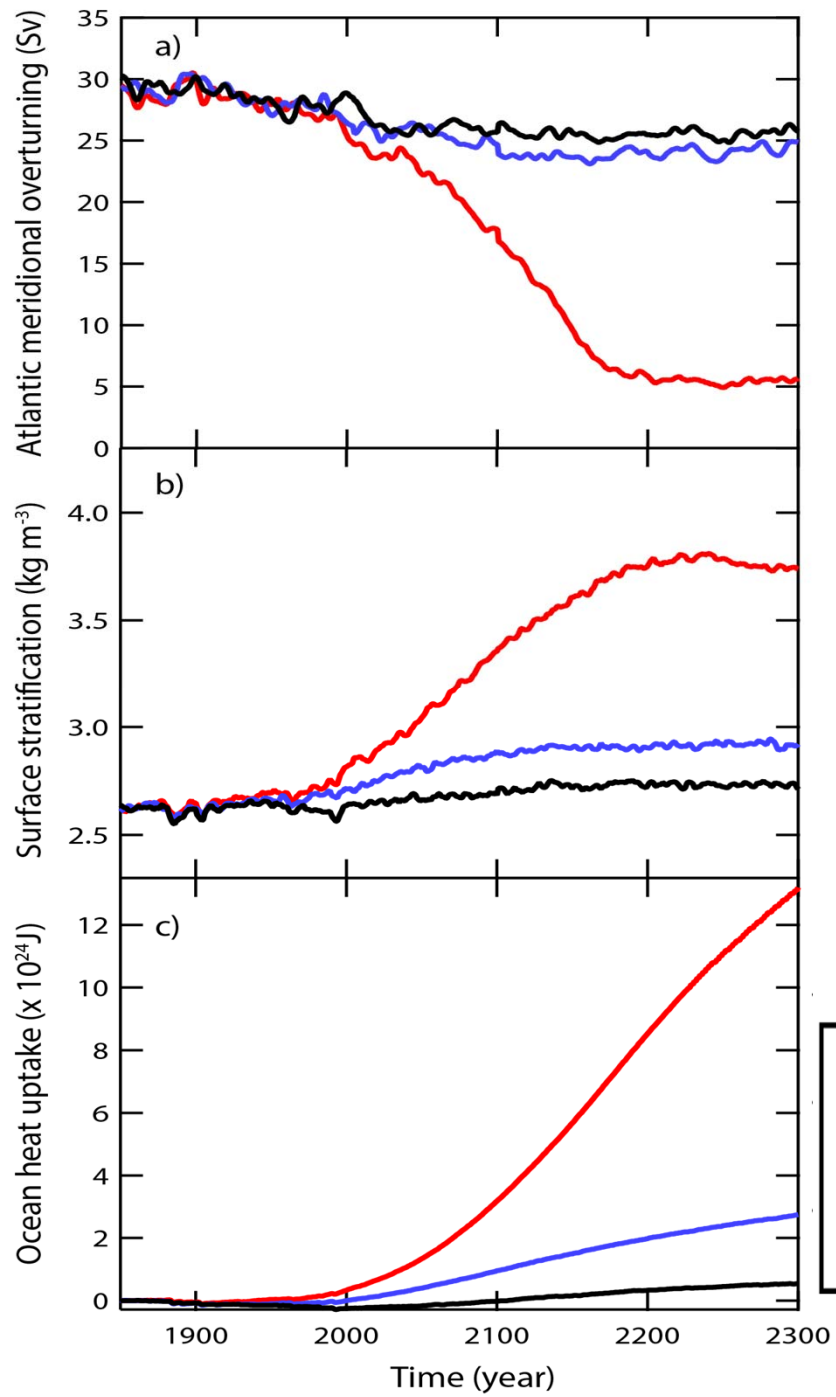
Parameter	Time Period			
	1850-1999	1850-2100	1850-2200	1850-2300
α (K/ppm)	0.0080	0.0048	0.0037	0.0041
β_L (Pg C/ppm)	-0.65	-0.18	-0.02	0.01
β_O (Pg C/ppm)	1.15	0.77	0.65	0.79
γ_L (Pg C/°C)	-2.9	-8.5	-16.4	-28.1
γ_O (Pg C/°C)	-1.5	-10.1	-24.4	-36.7
Gain (g)	0.013	0.034	0.056	0.091

$$g = \alpha(\gamma_O + \gamma_L) / (m + \beta_O + \beta_L)$$

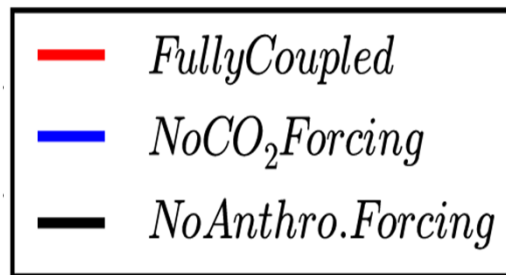
Ocean contributions to the climate-carbon feedback overtake land after 2100



Blue = FC – no CO_2 ; Red = FC – no anthro.; grey = no CO_2 – no anthro.

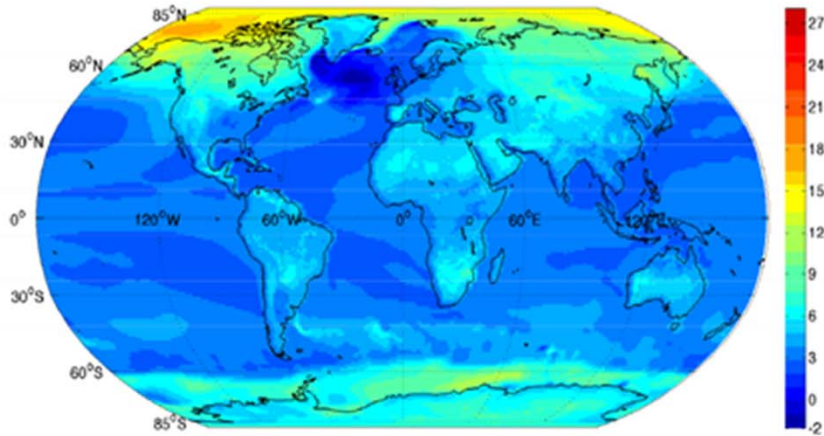


CESM1(BGC) temperature and salinity drivers of stratification at 2100

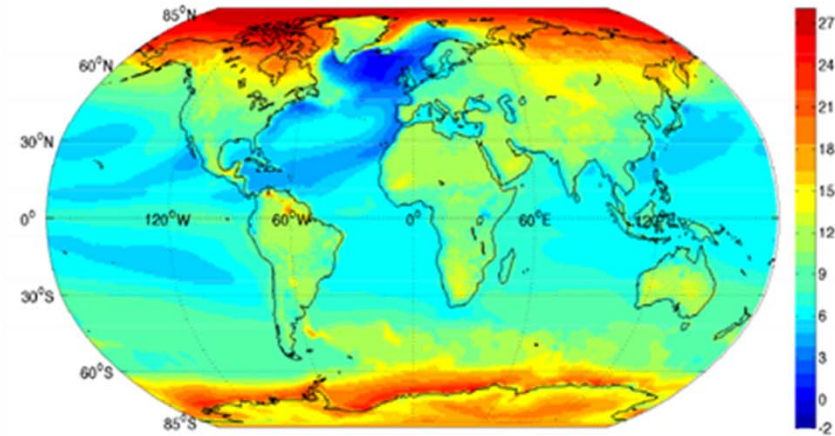


Shutdown in Atlantic Meridional Overturning Reduces Carbon Uptake in CESM

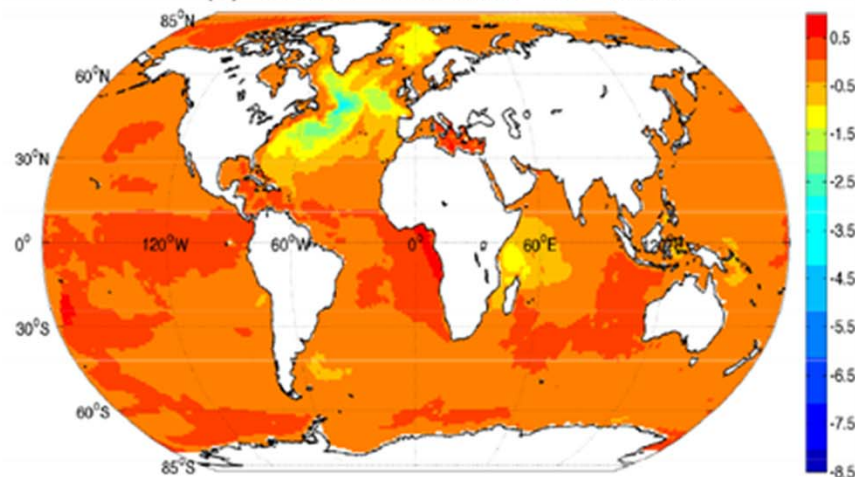
(a) T_{AS} : 2100-1850



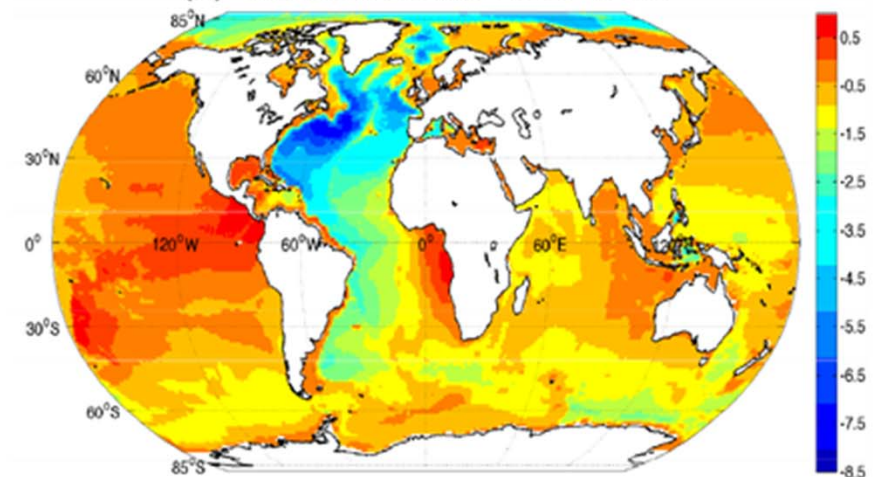
(b) T_{AS} : 2300-1850



(c) ocean carbon: 2100-1850 Kg C per m²

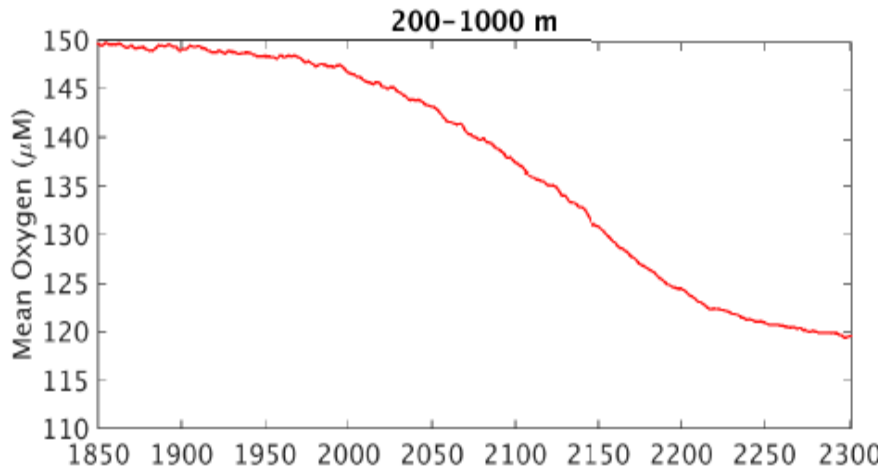


(d) ocean carbon: 2300-1850

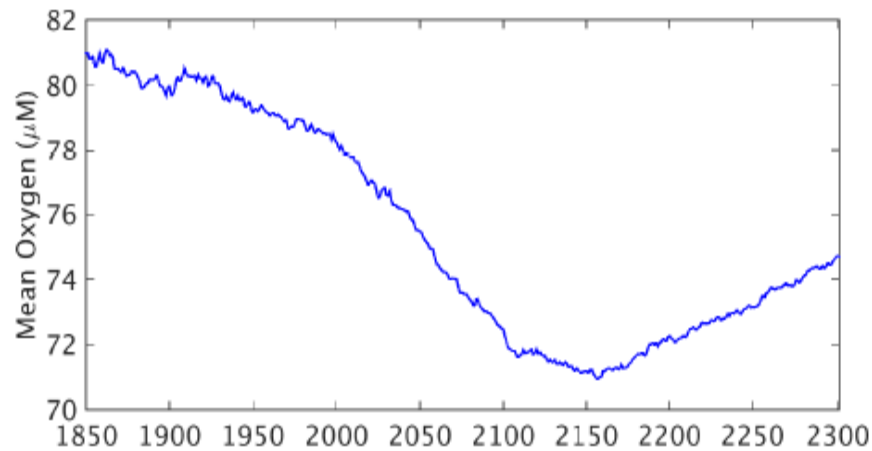


Randerson et al. (2015) GBC

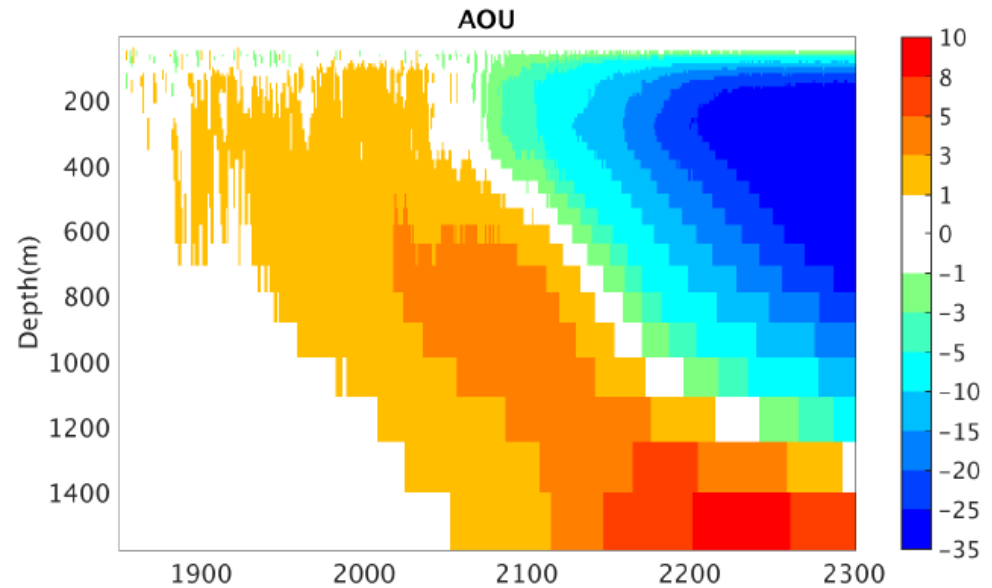
Tropical ocean anoxia reverses after 2100



(a) Global Mean O_2



(b) Mean O_2 in 30S-30N



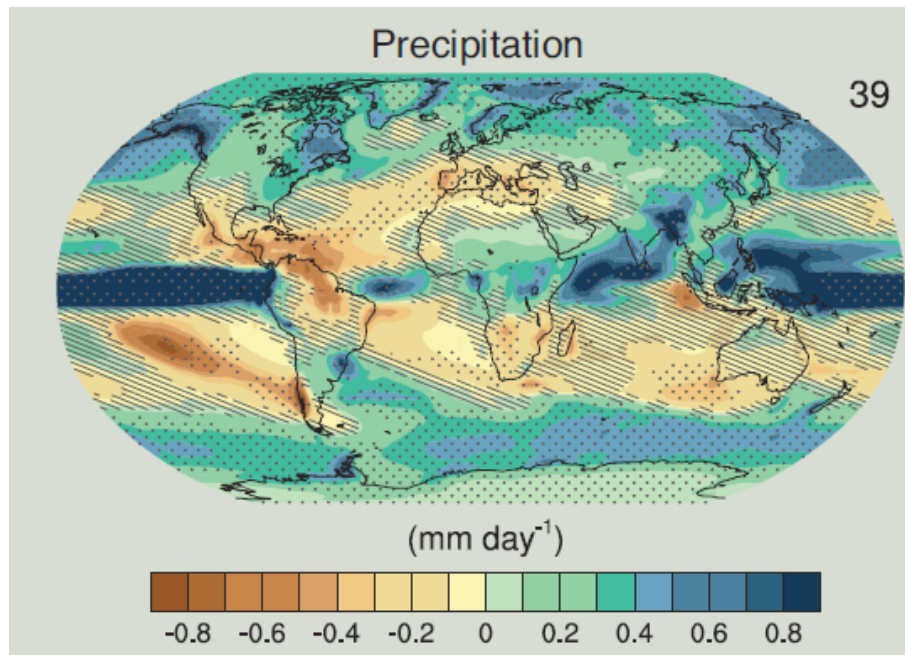
(b) AOU changes relative to 1850-60

Fu, Moore, Lindsay, Randerson, in prep.

Vulnerability of Central America and northern South America to changing drought

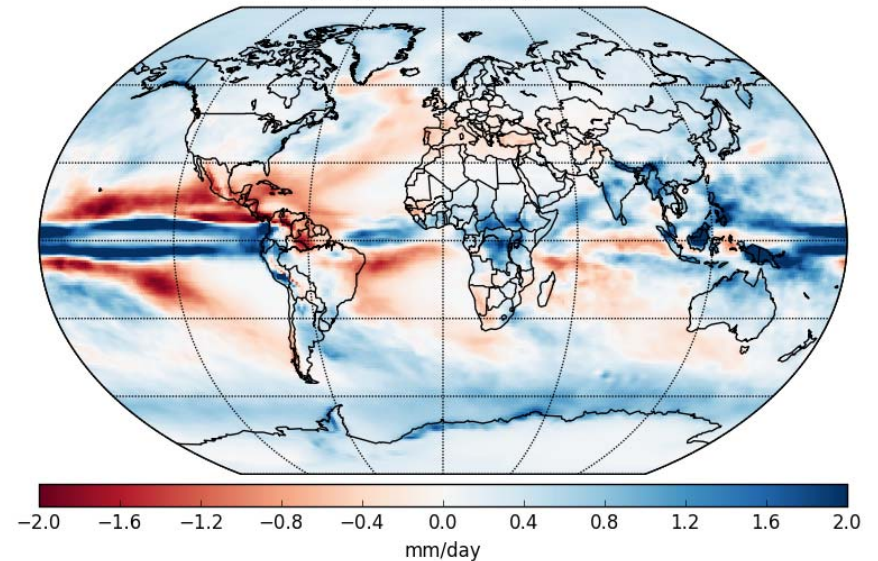
Precipitation changes for Representative Concentration Pathway 8.5
(2081-2100) – (1986-2005)

CMIP5 multi-model mean, IPCC AR1 TS



CESM1(BGC)

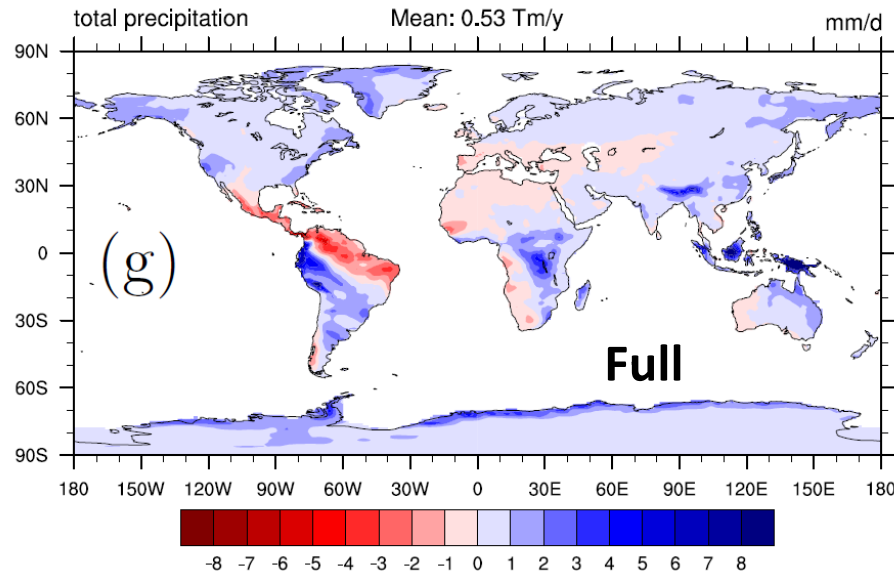
CESM1(BGC) Precipitation Difference Analysis: 1986-2005 to 2081-2100
RMSE: 0.608665791572



Hydrological cycle changes are not uniform across tropical land, with most models drying more in South America than in Africa or Asia

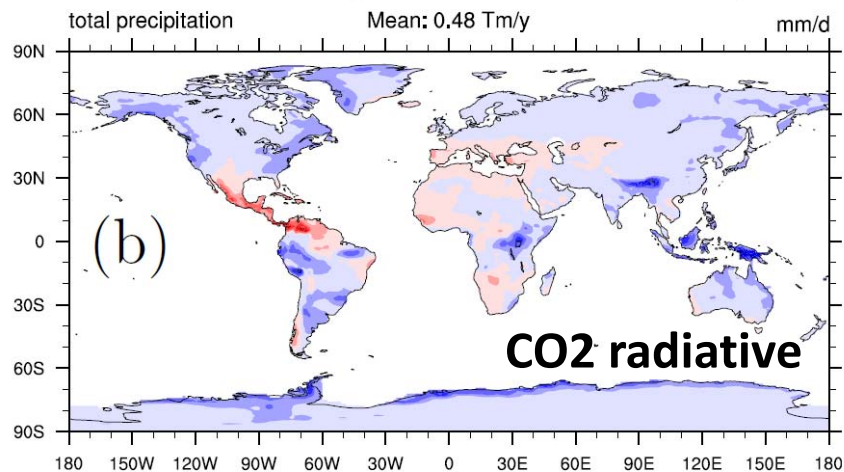
Precipitation reductions in neotropical forests driven equally by radiative and physiological effects of CO₂

$\Delta FC \text{ PRECIP}$ (2291–2300 minus 1851–1860)

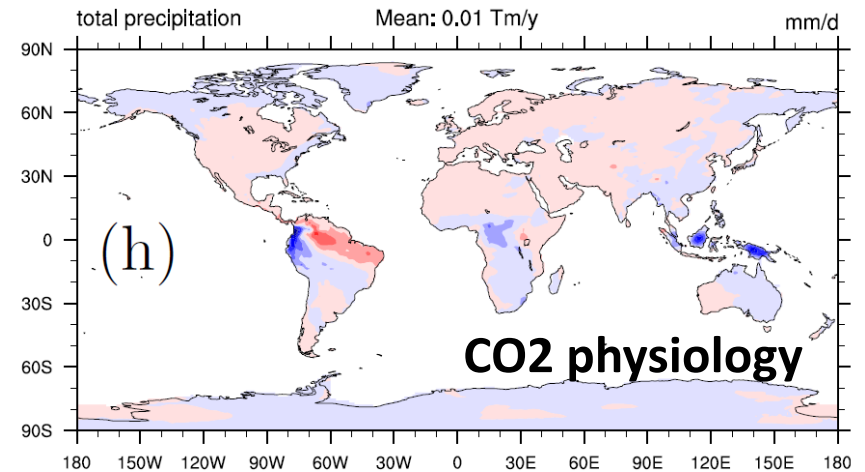


Hoffman (2015)
Ph.D. thesis, ms.
in prep.

$\Delta RAD \text{ PRECIP}$ (2291–2300 minus 1851–1860)

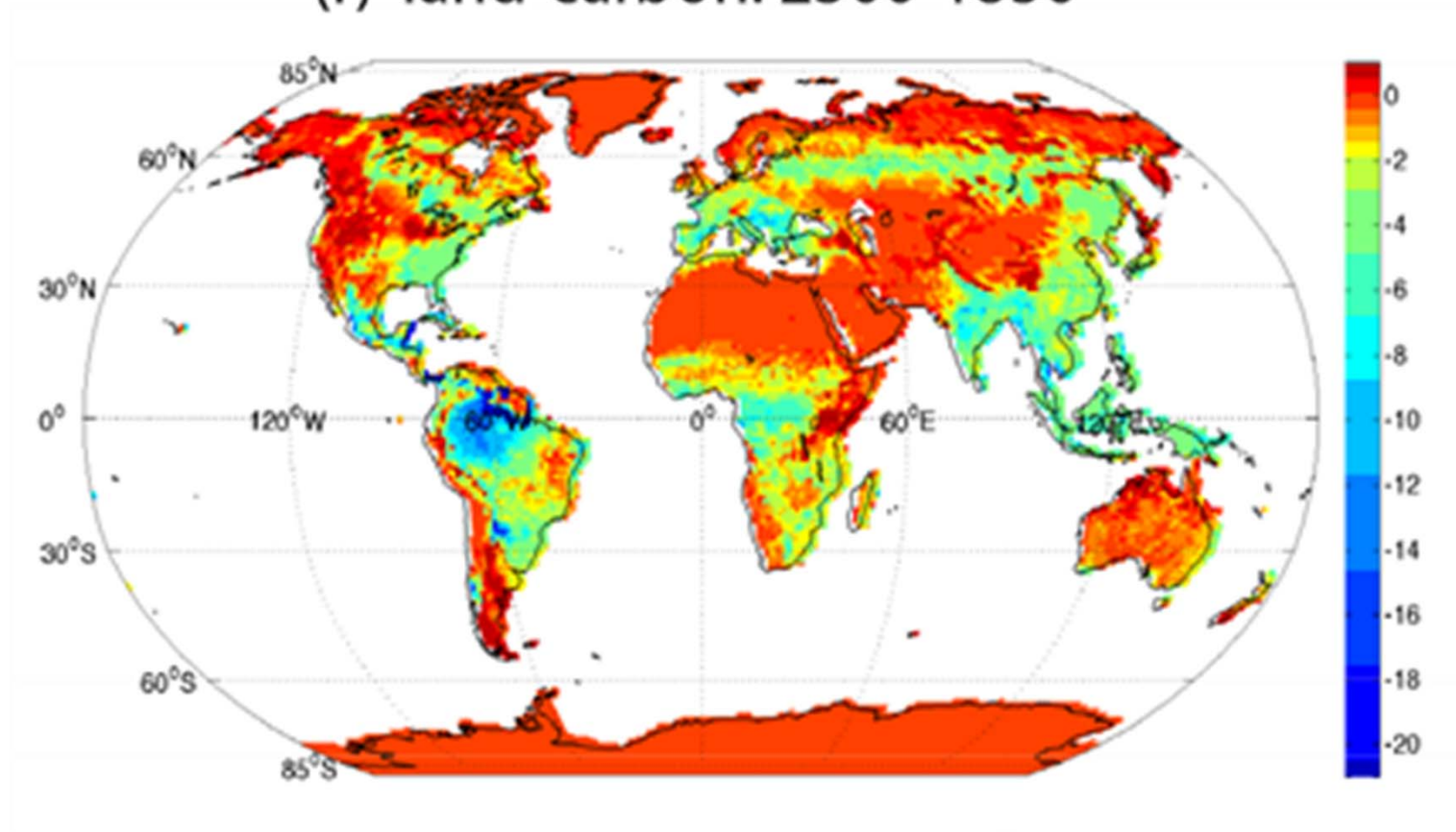


$\Delta BGC \text{ PRECIP}$ (2291–2300 minus 1851–1860)



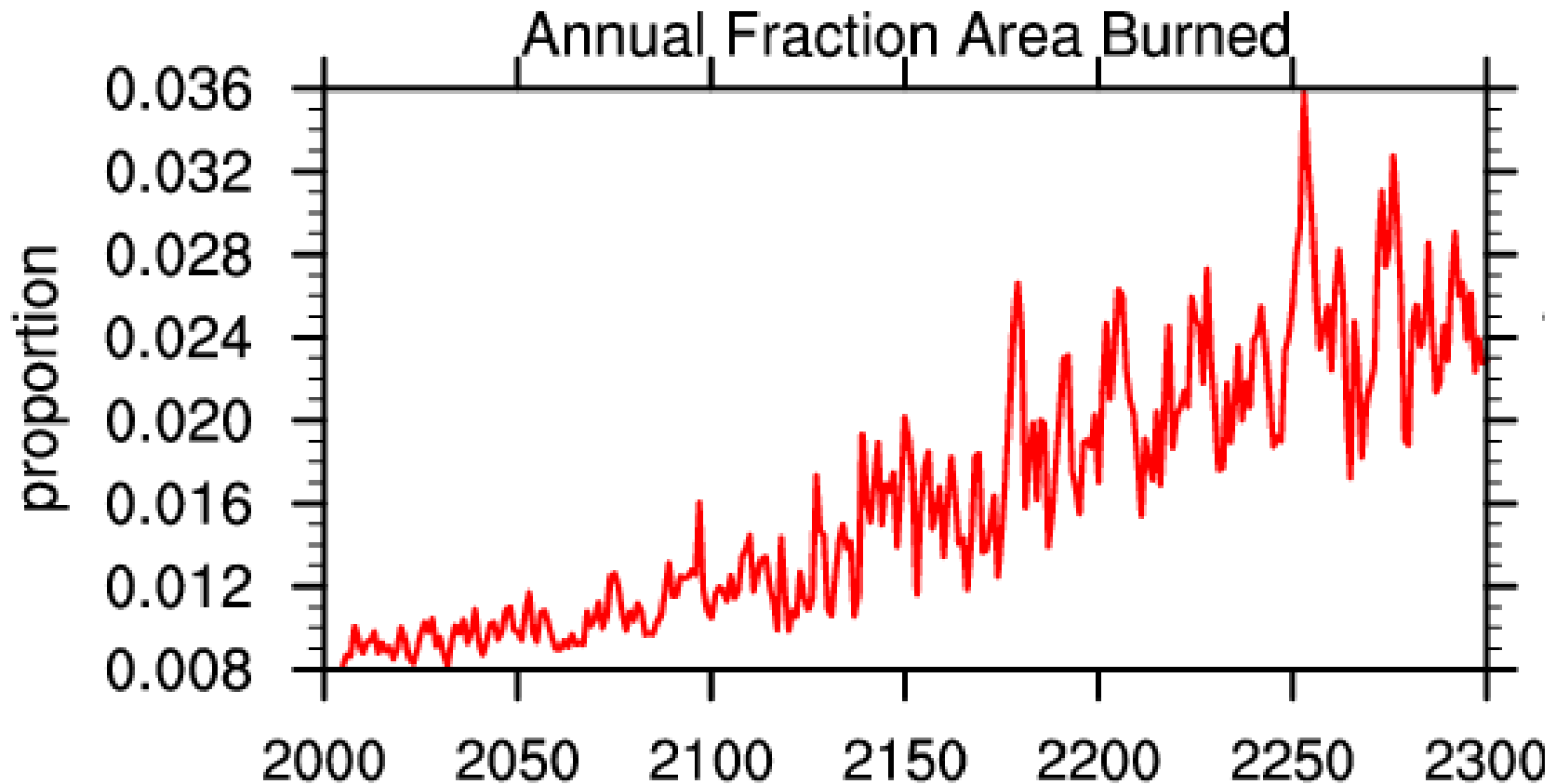
Forests in Central and South America exhibit a high degree of vulnerability to climate change-induced carbon loss

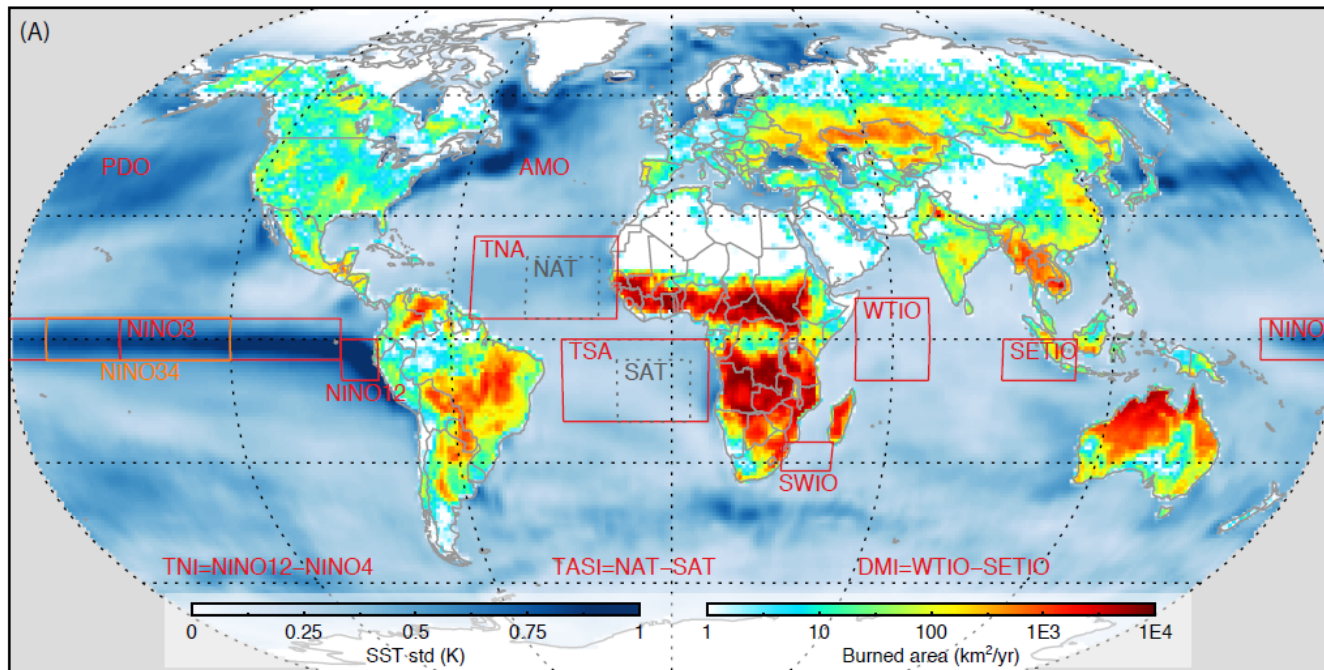
(f) land carbon: 2300-1850



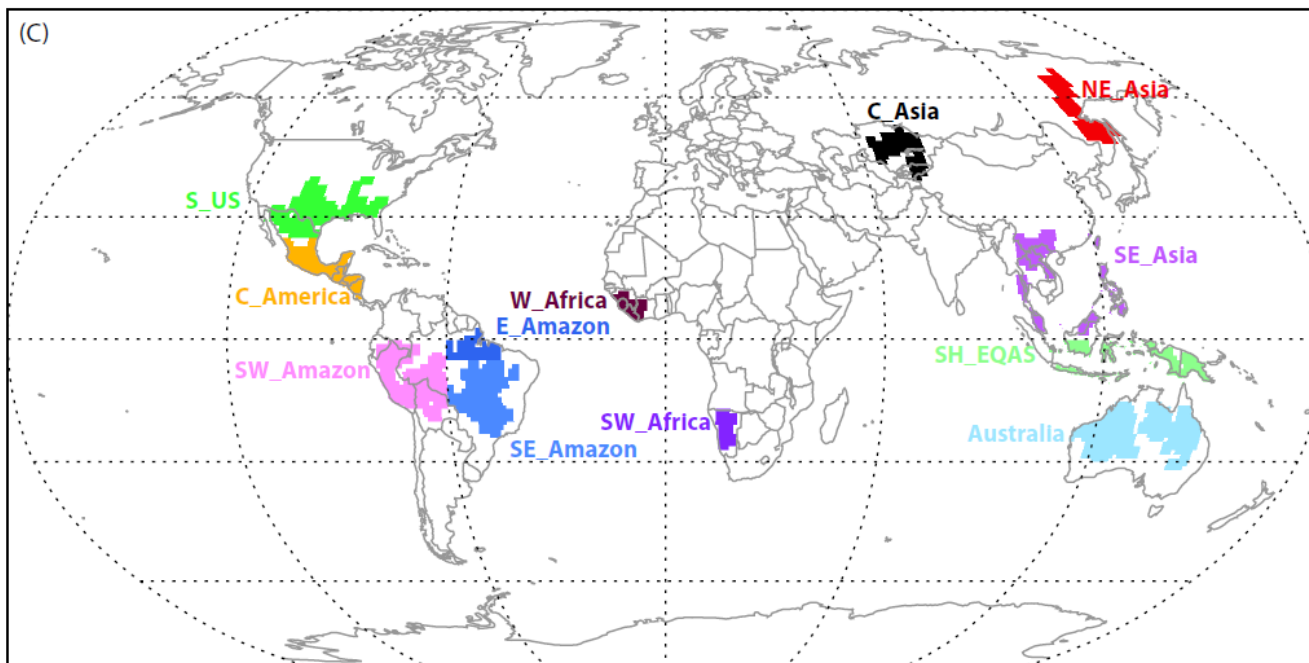
Kg C per m²

Amazon broadleaf forest burned area from the fully coupled simulation



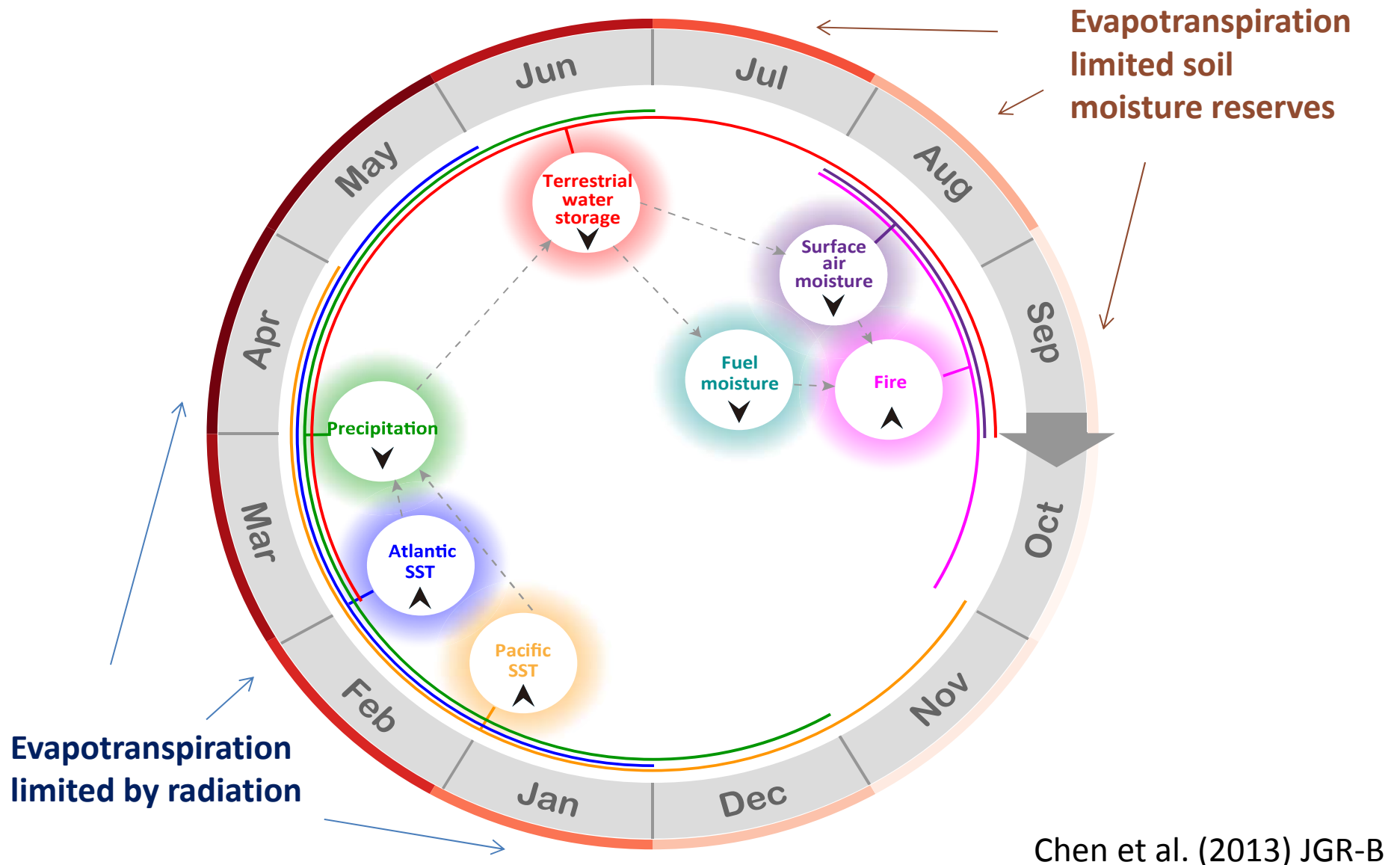


Toward the development of a global early warning system for fires

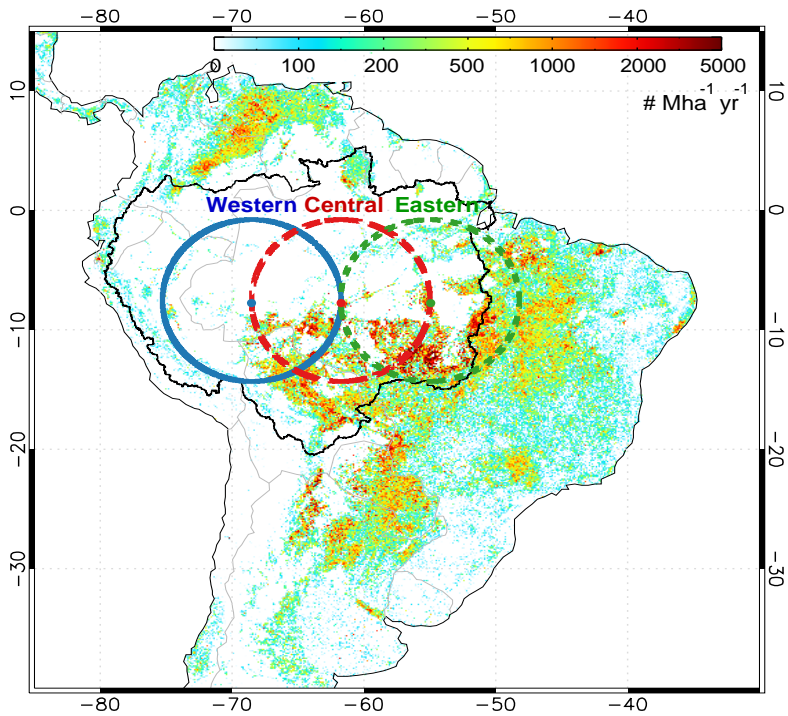


Chen,
Randerson, et
al. In prep.

A conceptual model for fire predictability in the Amazon is based on a forest soils capacitor mechanism



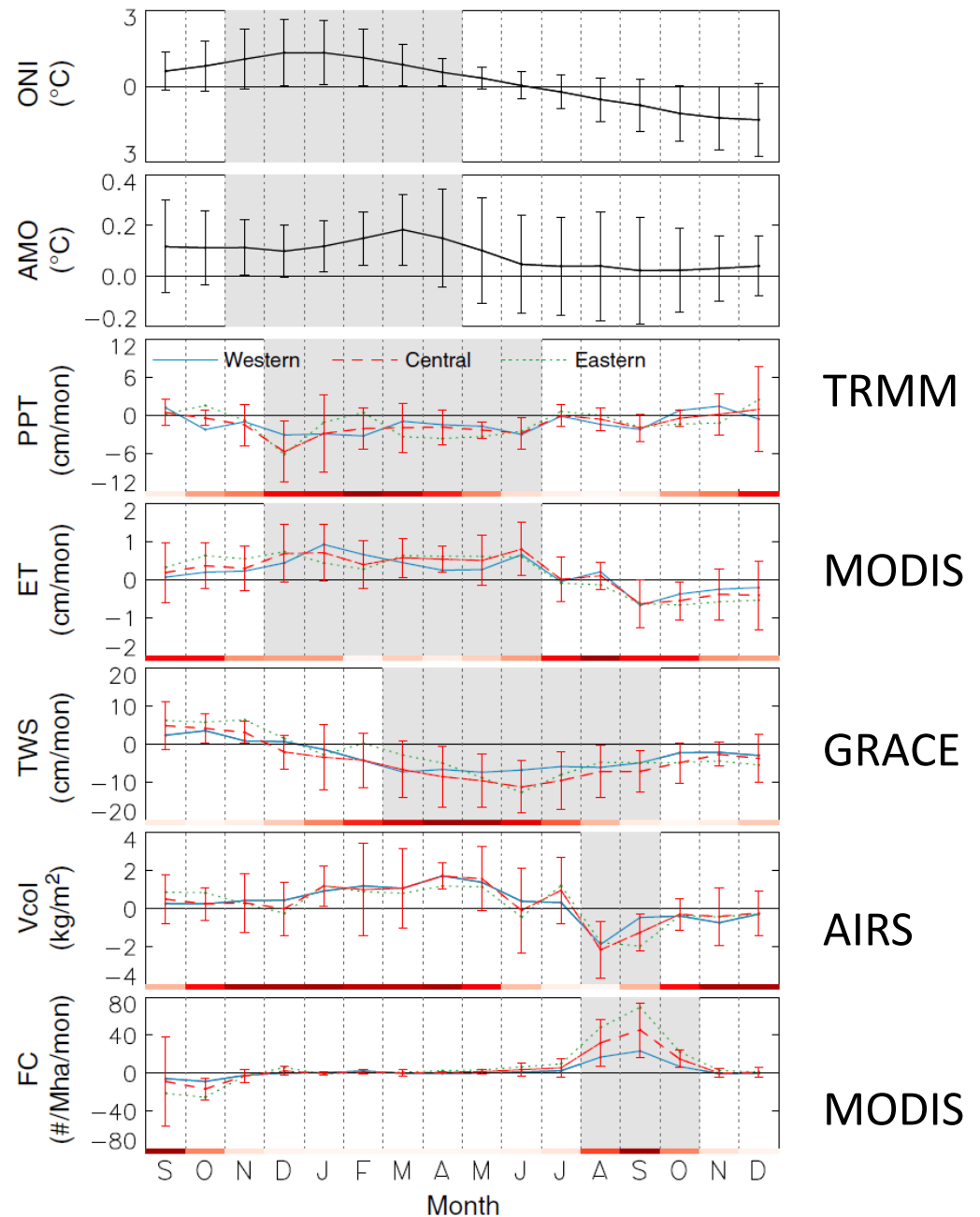
Test of the forest – soil moisture capacitor hypothesis for fire season predictions using satellite observations



High fire years: 2004, 2005, 2007, and 2010
 Low fire years: 2006, 2008, 2009, and 2011

Chen et al. (2013) JGR

Mean of high – low fire years



TRMM

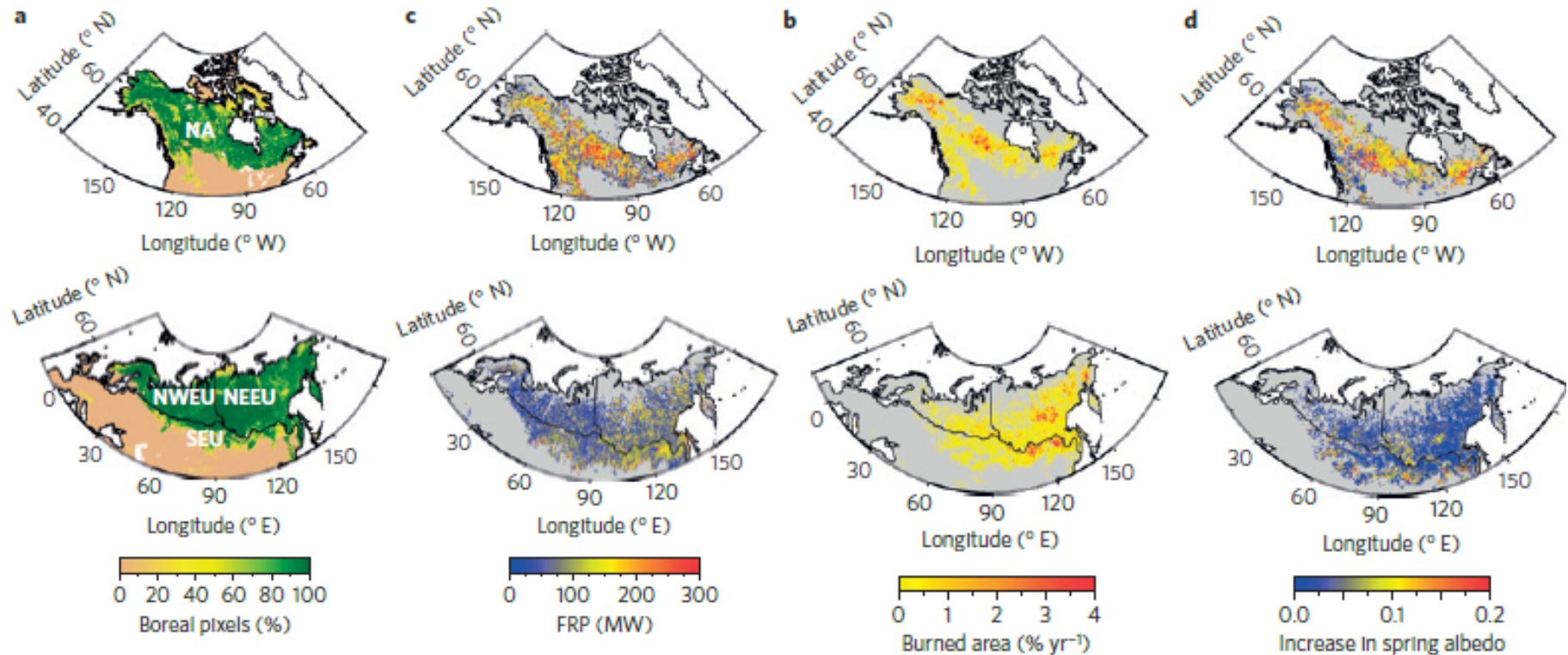
MODIS

GRACE

AIRS

MODIS

Capturing fire-mediated feedbacks in ESMs may require representing species-level effects



The presence of a single species in North America (black spruce) may cause fires to burn hotter and have greater long-term climate cooling effects

Rogers, Soja, Goulden, and Randerson. 2015. Influence of tree species on continental differences in boreal fires and climate feedbacks. *Nature Geosciences*. DOI: 10.1038/NGEO2352.

Conclusions

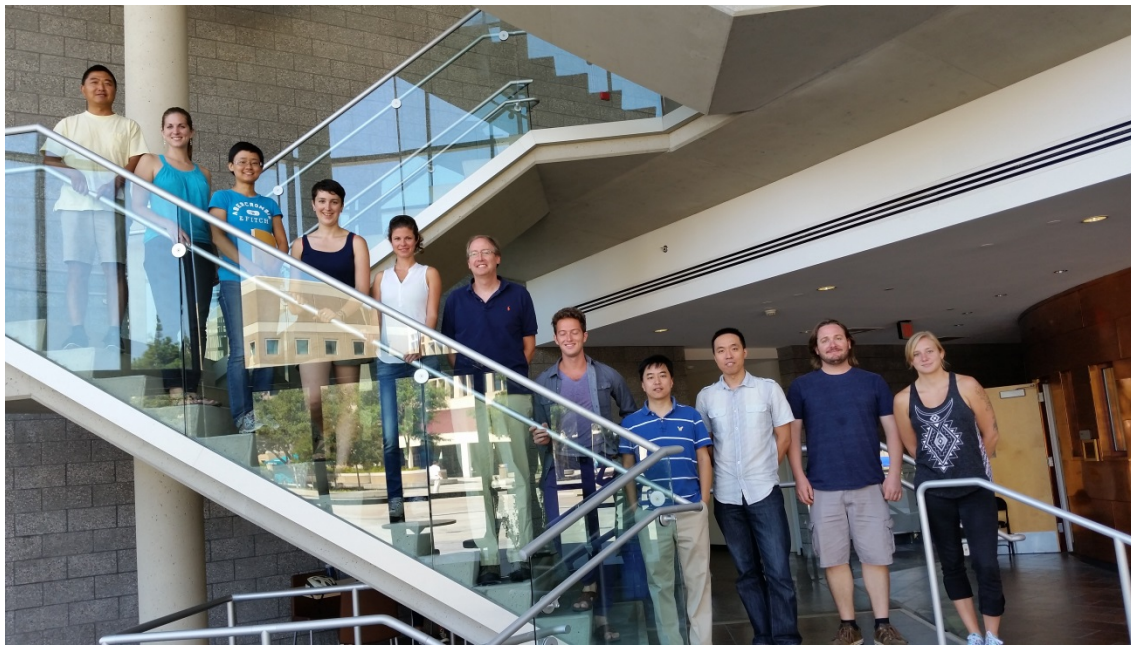
- Our understanding of Earth system dynamics, including processes that may contribute to ecosystem collapse, is woefully incomplete beyond 2100
- Ocean contribution to the climate-carbon feedback increases considerably over time for a “business as usual” scenario, and exceeds contributions from land after 2100
 - Land feedback likely reduced from land use change
 - Ocean feedback strength closely related to ocean heat content and AMOC shutdown
- Forcing from non-CO₂ agents for the RCP8.5 scenario is almost enough to surpass the 2 °C dangerous interference limit
- Tropical forests in Central and South America have a higher vulnerability to climate change than other tropical regions
- A better understanding and representation of fire processes in ESMs is essential for accurately predicting carbon cycle dynamics in drought-prone areas

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